



Metals content in surface waters of an upwelling system of the northern Humboldt Current (Mejillones Bay, Chile)

Jorge Valdés ^{a,*}, Domingo Román ^b, Gabriel Alvarez ^c, Luc Ortlieb ^d, Marcos Guíñez ^{a,e}

^a *Laboratorio de Sedimentología y Paleoambientes, Instituto de Investigaciones Oceanológicas, Facultad de Recursos del Mar, Universidad de Antofagasta, Casilla 170, Antofagasta, Chile*

^b *Departamento de Química, Facultad de Ciencias Básicas, Universidad de Antofagasta, Antofagasta, Chile*

^c *Departamento de Geomensura, Facultad de Ingeniería, Universidad de Antofagasta, Antofagasta, Chile*

^d *UR 055, PALEOTROPIQUE, Institut de Recherche pour le Développement, France*

^e *Programa de Doctorado en Ciencias Aplicadas, Facultad de Recursos del Mar, Universidad de Antofagasta, Chile*

Received 1 August 2005; received in revised form 26 April 2007; accepted 11 May 2007

Abstract

Physical–chemical parameters (temperature, salinity, dissolved oxygen, nutrients, and chlorophyll concentration) of surface waters were used to evaluate the influence of biological and physical processes over the metal concentrations (Cd, Ni, V, Mo, Mn, and Fe) in different periods of a normal annual cycle (June 2002 and April 2003), in Mejillones Bay (23° S), one of northern Chile's strongest upwelling cells. Two points were sampled every 2 months, but statistical analysis of these parameters did not show any spatial differences in surface water composition (annual average) in this bay. The order of total and dissolved metals by abundance (annual mean) in the Mejillones Bay surface waters during the sampling period was Cd < Ni < Mn < Fe < V < Mo.

The surface concentration of metals does not appear to be explained by anthropogenic inputs (at least not during the year of this work), and variability observed in this study appears to be natural. The lack of correlation between physical–chemical parameters and metals could indicate a more complex combination of factors acting on surface concentrations. Moreover, the mixture of water masses and the Oxygen Minimum Zone which characterizes the Mejillones bay should have an important influence on surface distribution of trace metals and can explain the high temporal variability observed in most of the metals analyzed in this work. A two-box conceptual model is proposed to suggest possible influences on metals in surface waters of this coastal ecosystem.

© 2007 Elsevier B.V. All rights reserved.

Keywords: Trace metals; Upwelling; Humboldt system; Marine surface water

1. Introduction

The continental margin is a boundary where several sources may affect the chemical characteristics of the

water and bring contaminating materials. Of these, the most important are continental sources (atmospheric transport, river runoff), marine (upwelling), and diagenetic exchanges at the water–sediment interface. These sources provide organic and inorganic materials that play an important role in biogeochemical cycles. For example, nutrients and trace metals are essential to the ecosystem, but should be carefully monitored because they are potential contaminants in the coastal environment (Cotte-Krief et al., 2000). In both cases, and in the

* Corresponding author. Tel.: +56 55 637865; fax: +56 55 637804.

E-mail addresses: jvaldes@uantof.cl (J. Valdés), droman@uantof.cl (D. Román), galvarez@uantof.cl (G. Alvarez), Luc.Ortlieb@bondy.ird.fr (L. Ortlieb), mguinez@uantof.cl (M. Guíñez).

42 absence of contamination processes, upwelling is the
43 principal source of surface nutrients and trace metals
44 from the deep ocean.

45 In general, all metals present in marine waters are in
46 dissolved and particulate forms. Many of these trace
47 metals are classified as micronutrients because they are
48 essential for phytoplankton growth. Most are compo-
49 nents of the enzyme system, which catalyzes important
50 biochemical reactions such as glycolysis, photosynthe-
51 sis, and protein metabolism. Partitioning between dis-
52 solved and particulate forms of trace metals depends on
53 the physical and chemical conditions of the marine
54 environment and the nature of each metal (Libes, 1992).
55 Moreover, any metal can assume various chemical
56 forms, being able to form a wide variety of ions, com-
57 pounds, or complexes, or being associated with different
58 mineralogical or organic phases (Grotti et al., 2001).
59 The concentration of these metals can be variable as a
60 consequence of changing inputs and/or seasonal effects
61 involving biological, geochemical, and physical inter-
62 actions (Hatje et al., 2001).

63 Knowledge of the biogeochemical cycle of trace
64 metals in coastal environments is needed in order to
65 identify pollution sources and to explore biological and
66 dynamical processes.

67 This paper aims to study the influence of biological
68 and physical processes over the some metals content in
69 different periods of a normal annual cycle, in surface
70 waters of one the most productive coastal systems of
71 northern Chile.

1.1. Regional setting

72

The Humboldt Current is one of the most productive 73
systems in the world. In the northern section of this 74
system, many permanent upwelling cells support im- 75
portant pelagic fisheries (Strub et al., 1998). Punta 76
Angamos and Mejillones Bay (Fig. 1) form the most 77
productive upwelling system of northern Chile. Many 78
studies concerning biological and physical characteris- 79
tics of this system have been developed in recent years, 80
improving our understanding about their response to 81
ocean-climatic variability, principally, El Niño events 82
(Escribano et al., 1998; Gonzalez et al., 1998, 2000; 83
Iriarte et al., 2000; Sobarzo and Figueroa, 2001; Ulloa 84
et al., 2001; Pizarro et al., 2002; Gonzalez et al., 2004; 85
Iriarte and Gonzalez, 2004). This area borders one of the 86
world's most arid regions, and continental input to the 87
ocean is restricted to minor atmospheric transport of 88
lithogenic particles (Vargas et al., 2004). For this reason, 89
upwelling seems to be one of the most important pro- 90
cesses influencing the chemical composition of the sur- 91
face waters. Mejillones Bay is a hotspot of biological 92
productivity. This bay is considered an upwelling 93
shadow system (UPS), which responds with more 94
intense biological productivity due to its internal thermal 95
stability (Marín and Olivares, 1999; Marín et al., 2003). 96
Studies of the Mejillones Bay have been, principally, 97
focused on the structure and composition of marine 98
sediments and their potential for paleo-ocean-climatic 99
reconstructions of the last millennia (Ortlieb et al., 2000; 100

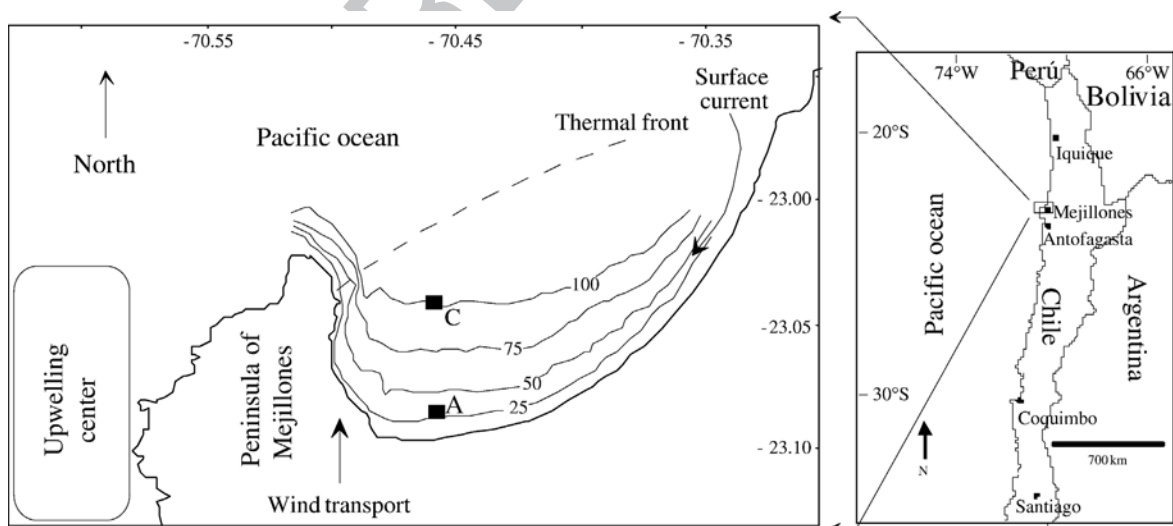


Fig. 1. Oceanographic setting of Mejillones Bay, showing schematic location of upwelling cell and location of sampling stations: A is coastal and C pelagic. Bathymetry's lines each 25 m are showed. The dashed line represents the position of the thermal front, which generates an upwelling shadow (UPS) in Mejillones Bay (Marín et al. 2003).

Valdés et al., 2000; Valdés and Ortlieb, 2001; Valdés et al., 2003, 2004; Vargas et al., 2004). As part of the latest studies, redox-sensitive metals preserved in sediments have been analyzed in order to calibrate proxies of paleoxygenation (Valdés, 2004) and to assess potential problems of marine contamination derived from intense industrial development in recent years (Valdés et al., 2005).

Industrial activity in the area is represented by aquaculture and the fishmeal industry, power thermoelectric generation, and mineral exports. The real impact of this recent development on the marine ecosystem is still unknown.

2. Materials and methods

2.1. Sample collection

Surface water samples were collected from two sites in Mejillones Bay (Fig. 1), using clean, 5 l inert sampling bottles (Niskin, General Oceanic) hung on stainless steel hydrowire. Between June 2002 and April 2003, samples for nutrients, chlorophyll-*a*, temperature, dissolved oxygen and salinity were collected monthly, while samples for trace metals were collected every 2 months. Samples were stored at low temperatures, in high density polyethylene bottles treated with nitric acid. Temperature, salinity, and dissolved oxygen data were collected with a CTD (Sea Bird, 19 plus) at the same sampling locations.

2.2. Analytical procedures

2.2.1. Trace metals

For dissolved concentrations, the samples were filtered with 0.45 μm Millipore membranes in Nalgene Polysulfone (PSF) systems and acidified to pH 2 with nitric acid. This was done the same day as the collection in a “clean laboratory environment”, inside a laminar flow hood (Labconco, Purifier Class II). The samples were maintained at 4 $^{\circ}\text{C}$ inside the hood. Immediately before the ICP–OES measurements, the samples were acidified to 1 M with HNO_3 .

For total concentrations, whole water samples were acidified to pH 2 with nitric acid, and then a 250 ml sub sample was digested with 20 ml of nitric acid at 90 $^{\circ}\text{C}$ in a BOD bottle for 2 h in a stainless steel air forced oven (WTC Binder); the final acidity of this medium was 1 M HNO_3 .

Fe and Mo were determined according to the method described by Román et al. (2003). Fe was measured with Hydraulic High Pressure Nebulization, Flame Furnace,

Atomic Absorption Spectrometry (HHPN–FF–AAS) before off-line separation and preconcentration with ammonium pyrrolidine dithiocarbamate (APDC) and extraction with methyl isobutyl ketone (MIBK). Mo was measured with Hydraulic High Pressure Nebulization, Atomic Absorption Spectrometry (HHPN–AAS) before off-line separation and preconcentration with MIBK; afterwards, the sample was acidified to 7 M with HCl.

V and Mn were determined by Inductively Coupled Plasma, Optical Emission Spectrometry (ICP–OES) according to the matrix matching technique, using purified seawater (PSW) (1) for blank measurements and the preparation of the standards for optimization and calibration.

Ni was determined by adsorptive accumulation of the respective complex with dimethylglyoxime at the drop mercury electrode (DME) by Adsorptive Differential Pulse Polarography (AdDPP) (2), and Cd was measured by Potentiometric Stripping Analysis (PSA) in the derivative signal mode (3, 4).

CASS — 4 (coastal seawater) and NASS — 4 (oceanic seawater) certified reference seawater from the Canadian National Research Council (NRC–CNRC) were used to validate the techniques and for quality control of the analytical determinations of Mn, Fe, Ni, Cd, and Mo. Vanadium’s quality control was done with the spiking technique. Purified coastal seawater (PSW) was prepared with an additional solid phase extraction step, using C_{18} 3 M Empore Bakerbond (J. T. Baker) discs (1) to improve the uptake of residual trace metals, and was used for blanks, the ICP–OES calibration matrix, with spikes as a secondary reference material for quality control.

Before each series of ten measurements routine controls of precision and accuracy were done. If, in four consecutive tests, the precision varied from the expected value by 20%, all analytical procedures were reviewed, and any having suspicious measurements were repeated. When the committed relative error was $\geq 15\%$ with respect to the certified value, all the analytical procedures were reviewed, and any suspicions ones were repeated.

All metal concentrations are expressed in $\mu\text{g l}^{-1}$.

2.2.2. Nutrients and chlorophyll

Water samples for nutrients were filtered through a 0.45 μm membrane filter and concentration was determined using a Spectronic 20D Spectrophotometer, according to methodology proposed by Parsons et al. (1984). Chlorophyll-*a* was measured in a Shimadzu RF-5301 Spectrofluorometer, after filtering samples in the dark through a 0.45 μm membrane filter, following

t1.1 Table 1
t1.2 Descriptive statistics of all parameters measured in this work

t1.3	Parameter	Mean		SD		Range		
		ST. A	ST. C	ST. A	ST. C	ST. A	ST. C	
t1.5	Temperature (°C)	17.02	17.04	1.79	1.56	14.47–19.92	15.18–20.06	
t1.6	Salinity (psu)	34.79	34.77	0.11	0.10	34.64–34.98	34.65–34.98	
t1.7	Dis. Oxygen (ml l ⁻¹)	6.40	7.24	2.31	2.02	2.22–10.47	3.20–10.68	
t1.8	Nitrate	2.13	1.04	2.95	1.15	0.06–8.70	0.02–3.66	
t1.9	Nitrite	0.11	0.08	0.14	0.07	0.00–0.46	0.00–0.19	
t1.10	Phosphate	1.40	1.13	0.74	0.49	0.20–2.70	0.30–1.18	
t1.11	Silicate	16.55	17.39	21.86	18.61	0.92–72.60	0.00–50.50	
t1.12	Chlorophyll- <i>a</i>	3.2	2.73	2.08	1.68	0.67–6.36	0.24–5.31	
t1.13	V	Total	7.33	7.36	1.91	3.21	4.40–9.23	3.48–12.80
t1.14		Dissolved	4.95	3.78	1.47	1.53	3.59–6.61	2.11–6.39
t1.15	Mn	Total	1.49	1.99	0.73	1.18	0.67–2.54	0.87–4.09
t1.16		Dissolved	0.94	1.03	0.31	0.33	0.56–1.32	0.56–1.38
t1.17	Fe	Total	2.77	2.69	0.41	0.67	2.41–3.51	1.96–3.58
t1.18		Dissolved	1.88	1.75	0.58	0.20	1.17–2.41	1.57–2.11
t1.19	Ni	Total	0.88	0.83	0.05	0.23	0.84–0.98	0.36–0.98
t1.20		Dissolved	0.57	0.57	0.23	0.25	0.20–0.88	0.21–0.82
t1.21	Mo	Total	13.85	15.28	1.79	5.02	11.19–15.36	9.43–21.74
t1.22		Dissolved	7.40	10.57	1.88	3.78	4.18–8.97	5.73–15.44
t1.23	Cd	Total	0.27	0.21	0.30	0.12	0.04–0.80	0.10–0.39
t1.24		Dissolved	0.04	0.04	0.04	0.03	0.01–0.11	0.01–0.07

t1.25 SD = standard deviation; nutrients are expressed in $\mu\text{mol l}^{-1}$; Chlorophyll is expressed in $\mu\text{g l}^{-1}$; metals are expressed in $\mu\text{g l}^{-1}$.

199 Holm-Hansen et al. (1965). The results of nutrients are
200 reported as $\mu\text{mol l}^{-1}$; chlorophyll-*a* is reported as $\mu\text{g l}^{-1}$.

3. Results and discussion

209

3.1. Statistical validation

210

2.3. Statistical analysis

202 Temperature, dissolved oxygen, salinity, metals,
203 nutrients, and chlorophyll were each analyzed by one-
204 way analyses of variance (ANOVA), to test the signif-
205 icance of spatial and temporal variability. Pearson corre-
206 lations and cluster analyses were performed to evaluate
207 the relationships between all parameters measured at
208 each station.

Descriptive statistics of all parameters measured in
211 both stations in Mejillones Bay are shown in Table 1.
212 The comparison of these results with previous work in
213 this area (Rodríguez et al., 1991; Marín and Olivares,
214 1999; Marín et al., 2003) indicates that the sampling
215 period corresponds to a normal year, without influence
216 of El Niño event, which drastically changes water
217 column structure in this area.
218

t2.1 Table 2
t2.2 Spatial variability: one-way ANOVA of all parameters measured in Mejillones Bay

t2.3	Temperature	Oxygen	Salinity	Chlorophyll- <i>a</i>	Nitrate	Nitrite	Phosphate		
t2.4	<i>F</i>	0.000	0.75	0.14	0.31	1.18	0.39		
t2.5	<i>P</i>	0.977	0.398	0.715	0.588	0.291	0.539		
t2.6	Silicate	V _{Total}	Mn _{Total}	Fe _{Total}	Ni _{Total}	Mo _{Total}	Cd _{Total}	V _{dissolved}	
t2.7	<i>F</i>	0.01	0.00	0.79	0.06	0.32	0.38	>0.26	1.83
t2.8	<i>P</i>	0.927	0.986	0.396	0.812	0.587	0.553	0.621	0.206
t2.9	Mn _{dissolved}	Fe _{dissolved}	Ni _{dissolved}	Mo _{dissolved}	Cd _{dissolved}				
t2.10	<i>F</i>	0.27	0.26	0.00	3.39	0.01			
t2.11	<i>P</i>	0.617	0.619	0.963	0.095	0.930			

t2.12 $P < 0.05$.

t3.1 Table 3
t3.2 Pearson's Coefficient Correlation of all parameters measured in Mejillones Bay

t3.3		Temperature	Oxygen	Chlorophyll	Nitrate	Nitrite	Mn _{tot}	Mo _{tot}	Cd _{tot}
t3.4	Oxygen	0.53							
t3.5	Chlorophyll	0.804							
t3.6	Nitrate	-0.601		-0.651					
t3.7	Nitrite				0.86				
t3.8	Phosphate		-0.7						
t3.9	Silicate		-0.7						
t3.10	Mn _{tot}			-0.639					
t3.11	Fe _{tot}	0.822							
t3.12	Ni _{tot}						0.652		
t3.13	Cd _{tot}	0.655							
t3.14	V _{dis}				0.67	0.715			
t3.15	Mn _{dis}			-0.825			0.832		
t3.16	Mo _{dis}			0.606				0.598	
t3.17	Cd _{dis}	0.656							0.699

t3.18 Only significant values ($P < 0.05$) are showed.

219 The results of the ANOVA showed no significant
220 differences for oceanographic parameters between the
221 two sampling stations (Table 2). In other words, the
222 surface waters of Mejillones Bay, at averaged over a
223 year, present homogeneous spatial conditions for all
224 parameters measured. However, the results suggest
225 some differences related to warm and cold periods
226 which are characteristics of this zone.

227 Table 3 presents the results of Pearson's Coefficient
228 of Correlation for all parameters measured in this study.
229 In spite of number of parameters measured in this study,
230 significant correlations were found in few cases. Only
231 temperature and chlorophyll correlate with more than
232 two other parameters.

3.2. Oceanographic variability

233

234 The annual mean surface temperature, salinity, and
235 dissolved oxygen levels found are characteristic of a
236 normal (non-El Niño) year, whereas the ranges of values
237 reflect seasonality and upwelling events influences.

238 Considering the results of the ANOVA test, cluster
239 analysis was done with all physical–chemical parameters
240 combining both sampling stations. The result (Fig. 2)
241 showed a clear group formed by temperature, oxygen, and
242 chlorophyll, which present positive and significant
243 correlations, except in case of oxygen and chlorophyll
244 relationship, but nitrate–nitrite (significant correlation,
245

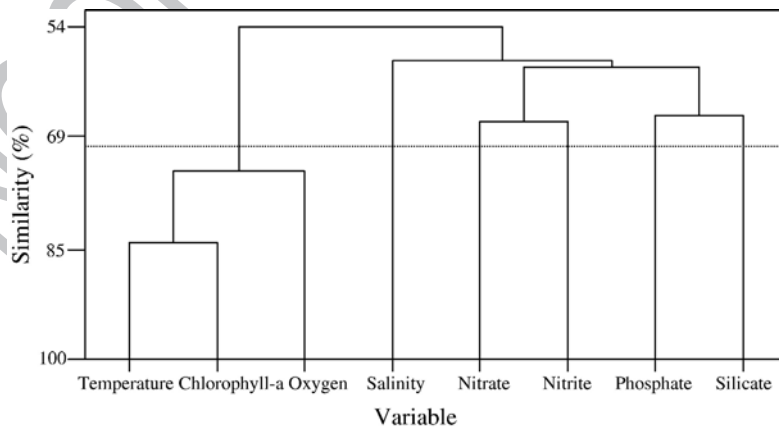


Fig. 2. Single-linkage dendrogram of similitude showing the results of clustering analyses of grouped physical–chemical parameters (not metals). The best line of similitude is shown.

Table 3), and phosphate–silicate (no significant correlation) suggest some relationship (Fig. 2). Supported in these results we interpret that Mejillones Bay system presents two oceanographic scenarios: one characterized by a low influence of physical processes and rich nutrient surface waters (cold season) and another characterized by strong influences of physical processes coupled to biological factors (warm season), the last one represented in Fig. 2 by group formed by temperature, oxygen, and chlorophyll.

Temperatures ranged between 14.5 °C (winter) and 20.1 °C (summer) and dissolved oxygen between 2.2 ml l⁻¹ and 10.7 ml l⁻¹ (both in summer; Table 1, Fig. 3). A considerable decrease of these parameters was registered in February. This situation suggests upwelled waters present inside the bay. At this month, it also registered an increase of phosphate (Fig. 3). The negative correlation between dissolved oxygen and phosphate has been used as an indicator of upwelling in this bay (Rodríguez et al., 1991).

During cold season (April–September) nutrients concentration is high because productivity is inhibited due to a lack of radiation. The low intensity of productivity generates low chlorophyll concentrations (Fig. 3). During warm season (October–March), when radiation is optimum, the productivity increases and, consequently, the nutrient concentrations decrease and chlorophyll increases (Fig. 3). A similar pattern, especially between nitrate and chlorophyll, was observed by Rodríguez et al. (1991), during 1987–88.

3.3. Metal content in surface waters

The concentrations of total and dissolved metals at each station are showed in Table 1. As was the case for oceanographic parameters, ANOVA did not show significant differences in metal concentrations between both sampling stations (Table 2). In order of abundance, the total and dissolved trace metals (annual mean) in the surface waters of Mejillones Bay (2002–2003) were Cd < Ni < Mn < Fe < V < Mo.

In general, metals in seawaters can be classified according to their interactive properties with biological and physical–chemical processes. Two groups can be identified: conservative and non-conservative. The first group, including Mo, presents few interactions with the biological cycle and its horizontal and vertical distribution is governed by physical processes such as advection and turbulent mixing (Libes, 1992). The second group, including V, Ni, Cd, Fe, and Mn, is affected by biological processes, whether *via* scavenging and/or biological uptake (Brown et al., 1994). In the latter

case, phytoplankton affects the distribution of dissolved metals, using most of them as micronutrients (Gonzalez-Davila, 1995). These intrinsic characteristics of metals must be considered when interpreting this study's results.

3.3.1. Vanadium

The predominant chemical form of V in oxic waters is vanadate, H₂VO₄⁻ (Morford and Emerson, 1999), whose behavior is mostly conservative with some possible nutrient cycling (Bruland, 1983). Total and dissolved V variability presents a general trend characterized by high concentrations in cold season and low concentrations in warm season, similar to nutrient behavior. The partitioning between the dissolved and particulate phases appears to be relatively constant throughout the year with exception of February in pelagic station (Fig. 3). This situation points to an increase of the particulate form, which could be associated to the influence of upwelled water transported into the bay by onshore advection.

3.3.2. Manganese

Mn is present in oxic waters as MnO₂ (Russell and Morford, 2001). According to Kremling (1985) and Libes (1992), this metal is enriched in coastal surface waters due to a large metal supply from external sources. Variability was more marked for total Mn than for its dissolved form, which could be explained by the presence of particulate Mn in surface waters during periods of high total Mn concentrations (*i.e.* October). According to Schenau et al. (2002), the most important primary source of Mn to the ocean is the product of continental weathering. Valdés et al. (2005) found a significant correlation between this metal and Al in the Mejillones Bay marine sediments, which is a signal of lithogenic input. In Mejillones the continental transport to the ocean is controlled almost exclusively by winds (Vargas et al., 2004). The strongest winds from the south and south-west develops during spring and summer conditions (Escribano et al., 2004) mostly in September–October, which could to explain the increase of particulate Mn registered in October.

3.3.3. Iron

The annual variation of Fe was different at each sampling station. At the coastal station, both total and dissolved iron concentrations were low in winter and increased during spring–summer. On the other hand, the pelagic station showed an opposite trend: high total concentrations during winter and low total concentrations during spring–summer, with constant dissolved

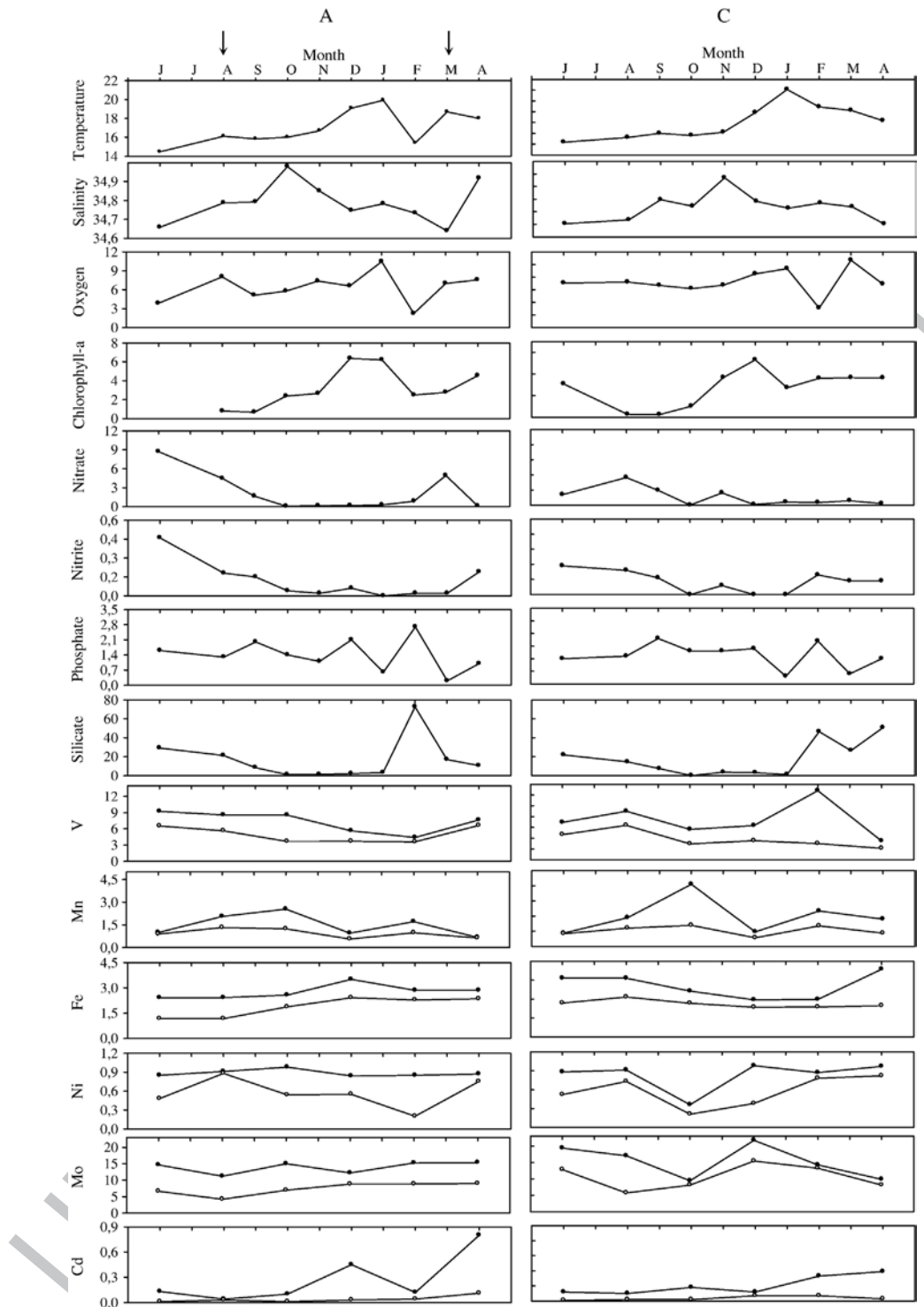


Fig. 3. Surface temporal variability of parameters measured at both Mejillones Bay stations during 2002–2003. Temperature in degrees Celsius; dissolved oxygen in ml l^{-1} ; chlorophyll-*a* in $\mu\text{g l}^{-1}$; nutrients in $\mu\text{mol l}^{-1}$; metals in $\mu\text{g l}^{-1}$. Black circle is total metal and white circle is dissolved metal. Arrows indicates satellite image dates (See Fig. 4).

345 concentrations year-round (Fig. 3). Martín and Fitzwater
 346 (1988) have shown that phytoplankton growth is limited
 347 by iron availability, especially in the California and
 348 Humboldt Current Systems (Hutchins and Bruland,
 349 1998). However, the lack of correlation between Fe and
 350 nutrients and chlorophyll would indicate that this
 351 metal has no relationship with biological factors, at least
 352 in this bay. By other hand, Des Combes et al. (1999)
 353 indicate that the supply of this metal is governed,
 354 principally, by continental input probably in the same
 355 way as Mn in Mejillones Bay. However, no significant
 356 correlation was found between Fe and Mn (Table 2). In
 357 the California Current Johnson et al. (1999) found
 358 that Fe was introduced by resuspended bottom sediment
 359 during coastal upwelling. A similar mechanism should
 360 be acting in Mejillones bay, because the microoxic
 361 condition of bottom environments (Valdés, 2004) and the
 362 persistence of the upwelling should to introduce Fe into
 363 the bay from offshore zone.

364 3.3.4. Nickel

365 Ni occurs in oxic waters as Ni^{2+} and NiCl^+ (Calvert
 366 and Pedersen, 1993). At the Mejillones Bay coastal
 367 station, total Ni was constant during the entire study
 368 period, whereas its dissolved form showed high
 369 concentrations during winter and low concentrations
 370 during spring–summer. The pelagic station registered a
 371 highly variable Ni concentrations, with decay during
 372 October and high concentrations before and after this
 373 month (Fig. 3). The distribution of Ni is strongly
 374 mediated biologically (Kremling, 1985), and its behav-
 375 ior as a micronutrient indicates that it is removed from
 376 surface waters by plankton growth (Saager et al., 1992;
 377 Morley et al., 1993). It seems like lowest dissolved Ni
 378 concentration at station A should be accompanied by a
 379 high chlorophyll concentration (Fig. 3), which sug-
 380 gest some influence of biological productivity on Ni
 381 behavior.

382 3.3.5. Molybdenum

383 Molybdenum (VI) appears in the stable oxidation
 384 state in oxic waters and is found as MoO_4^{2-} (Morford
 385 and Emerson, 1999; Russell and Morford, 2001;
 386 Nameroff et al., 2002). Mo is a conservative metal
 387 only affected by physical–chemical factors (Libes,
 388 1992; Brown et al., 1994; Adelson et al., 2001).
 389 Because of its conservative nature, it is possible to
 390 assume that coastal surface circulation controls the Mo
 391 distribution in Mejillones Bay surface waters. Concen-
 392 trations of dissolved and total Mo were less variable at
 393 the coastal than at the pelagic station (Fig. 3), which is
 394 closer than upwelling cell (Fig. 1). During warm season

the upwelling events are more frequent and intense 395
 (Marín et al., 2003), which influence the onshore 396
 advection and probably the Mo content in surface 397
 waters principally in the pelagic station. 398

399 3.3.6. Cadmium

Of all metals measured in this study, Cd had the 400
 lowest concentration. The total concentration presented 401
 a more marked variability than its dissolved form at both 402
 stations (Fig. 3), as a consequence of the variation of 403
 particulate Cd in Mejillones Bay surface waters. The 404
 typical oceanic Cd distribution presents low or depleted 405
 values near the surface (Delgadillo-Hinojosa et al., 406
 2001) and is regulated by marine biogeochemical pro- 407
 cesses, namely uptake by phytoplankton in surface 408
 waters (Abe, 2001). Cd is a nutrient-like metal that 409
 presents only one oxidation state in seawater (Cd II) and 410
 is found as CdCl^+ (Morford and Emerson, 1999; Russell 411
 and Morford, 2001; Nameroff et al., 2002). This 412
 behavior indicates that Cd has a short residence in 413
 surface waters because it is removed rapidly by plankton 414
 growth (Calvert and Pedersen, 1993). Temporal Cd 415
 variations were associated with chlorophyll-*a* concen- 416
 tration and temperature fluctuation, suggesting that 417
 phytoplanktonic biomass production is the principal 418
 factor controlling Cadmium concentration in surface 419
 waters of Mejillones bay. However, the two highest total 420
 Cd values measured in December and April in station A 421
 seem too high to be attributed to phytoplankton, because 422
 according to Cullen and Sherrell (1999) in a work 423
 developed in California Current, the Cd content of 424
 marine particles is less than 10 ng l^{-1} . It is possible that 425
 upwelled waters transported by onshore advection have 426
 some degree of influence in Cd content in Mejillones 427
 bay, as account in others coastal environments (Boyle, 428
 1988; Van Geen et al., 1992; Van Geen and Husby, 429
 1996; Takesue et al., 2004). 430

431 3.4. Factors controlling seasonal variability of trace 432 metals in Mejillones bay

One possible source of trace metals in Mejillones 433
 Bay is the input of lithogenic material. However, Valdés 434
 et al. (2000) and Valdés (2004) compared V, Ni, Mo, and 435
 Cd with aluminum (a marker of lithogenic sources 436
 according to Dean et al. (1997) in marine sediments, 437
 without finding any significant correlations). The 438
 authors concluded that continental input in the sedi- 439
 ments of Mejillones Bay is not the principal source of 440
 these trace metals. Moreover, Vargas et al. (2004) in- 441
 dicated that lithogenic debris in the bottom marine 442
 sediments of Mejillones Bay represents less than 5% of 443

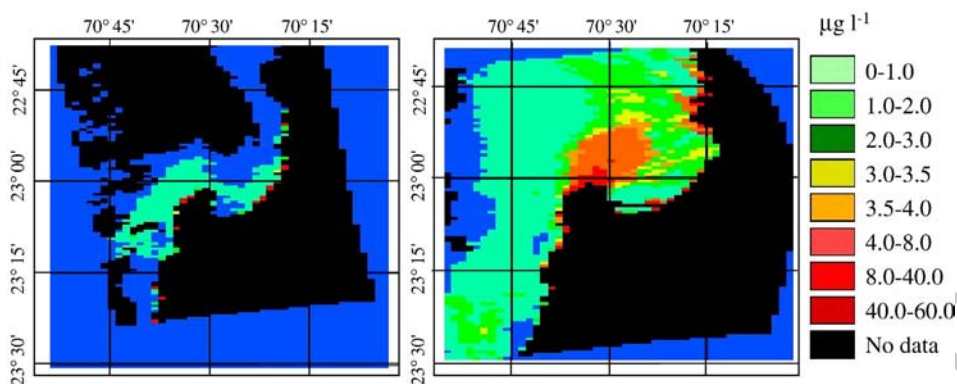


Fig. 4. Satellite image of chlorophyll-*a* in Punta Angamos upwelling system during August 2002 (left) and March 2003 (right).

444 the total bulk sediments, which suggests that metals in
 445 lithogenic minerals transported from the continent to the
 446 ocean constitute a minor source in this coastal area. By
 447 other hand, Valdés et al. (2005) demonstrated that
 448 contamination doesn't explain the concentration of these
 449 trace metals in this bay. However, some influence of
 450 waste water treatment plant effluent should be consid-
 451 ered specially in station A. Even if influenced to some
 452 extent by atmospheric inputs and/or contamination
 453 processes in the surface waters, the distribution of V,
 454 Ni, Cd, and Mo in Mejillones Bay at the moment
 455 of sampling, can be broadly and principally accounted
 456 by a combination of natural oceanic processes and
 457 mechanisms.

458 Seasonal variability was described by grouping param-
 459 eters into two general annual periods: the cold
 460 (autumn–winter) and warm (spring–summer) seasons.
 461 Fig. 4 shows typical satellite images (from EOS Aqua
 462 Spacecraft; second and third order polynomic correction
 463 process) of chlorophyll-*a* distribution during the cold
 464 period (August 28, 2002) and the warm period
 465 (February 28, 2003), coincident with the sampling of
 466 metals in surface waters. During the cold season low
 467 levels of chlorophyll were found in the surface waters of
 468 the Punta Angamos upwelling cell and Mejillones Bay
 469 (in accordance with chlorophyll measured *in situ*),
 470 which indicated that productivity was low during this
 471 period. On the other hand, during the warm season, well
 472 productive waters are common, with a filament of high
 473 levels of chlorophyll displaced to the north-east, gener-
 474 ating an upwelling shadow system (UPS) in the bay, as
 475 proposed by Marin et al. (2003). Low levels of dissolved
 476 oxygen and high levels of phosphate registered during
 477 warm season, particularly in February, are characteristic
 478 of upwelling events (Fahrbach et al., 1980).

479 The lack of a similar pattern of temporal variability of
 480 metals in both sampling stations seems to be the most

important characteristic. This situation could indicate 481
 that there is a combination of factors with different 482
 temporal variation patterns which influence the presence 483
 of metals in surface waters. Based on the area's oceanic 484
 characteristics and the results obtained in this study, a 485
 conceptual two-box model is proposed to explain the 486
 variation of metals in surface waters of this coastal 487
 ecosystem (Fig. 5). The two-box model considered in 488
 this work was based on the area's UPS interpretation 489
 proposed by Marin et al. (2003), which separates the 490
 Mejillones Bay from the Punta Angamos upwelling 491
 system. The upwelling event of Punta Angamos causes 492
 a biological productivity increment into Mejillones bay, 493
 and the development of an efficient retention zone of 494
 phytoplanktonic organisms (Marin and Olivares, 1999; 495
 Escribano et al., 2000). These characteristics are 496
 modulated by the “upwelling shadow” condition present 497
 in this bay, which is defined as a confined zone of an 498

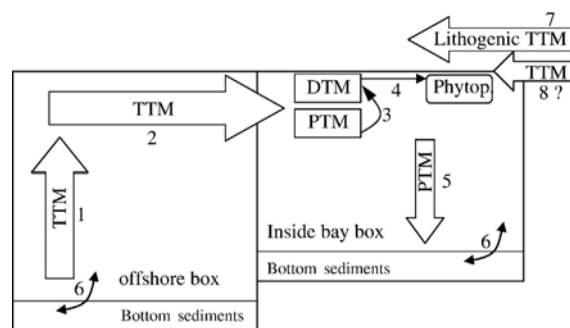


Fig. 5. Conceptual scheme of mechanisms which influence temporal variability of trace metal concentrations in surface waters of Mejillones Bay. TTM = Total Trace Metal; PTM = Particulate Trace Metal; DTM = Dissolved Trace Metal; 1 = offshore deep water upwelling; 2 = onshore advection; 3 = desorption of particulate metals; 4 = phytoplankton uptake; 5 = scavenging; 6 = exchange at sediment–water interface 7 = wind transport; 8 = expected anthropogenic input.

499 active coastal upwelling system, whiting which upwell- 551
500 ing is less intense (Marín et al., 2001). According to this 552
501 situation the upwelled waters are transported inside the 553
502 bay by onshore advection. The upwelling condition 554
503 founded inside the bay in February, according to PO4 555
504 and dissolved oxygen (Fig. 3) represents the spin-up 556
505 phase of the develop of the upwelling shadow (Marín 557
506 et al., 2001).

507 Several studies have shown that coastal surface 558
508 waters contain higher concentrations of dissolved trace 559
509 metals (DTM) than open ocean waters (Kremling and 560
510 Hydes, 1988; Kremling and Pohl, 1989; Muller et al., 561
511 1994; Cotté, 1997; Le Gall et al., 1999). The offshore 562
512 box presents two ways of metal transport: surface 563
513 circulation entering and vertical circulation *via* the 564
514 upwelling of deep water, which results in different 565
515 intensities depending on the wind direction and strength 566
516 (Strub et al., 1998). Both mechanisms transport water 567
517 masses with a much defined chemical composition 568
518 (metals included). For example, in the case of DTM, at a 569
519 global scale, Martín and Thomas (1994) estimated that 570
520 the oceanic input would dominate sources for the 571
521 oceanic margin. Unfortunately, information about the 572
522 metal composition of marine waters in the Humboldt 573
523 system is lacking and it was not possible to compare our 574
524 results with other works. In the inside the bay box, 575
525 surface metal concentrations result from very low 576
526 lithogenic input, uncertain industrial contamination pro- 577
527 cesses (a continuing possibility), ingestion by phyto- 578
528 plankton, desorption from particles, and scavenging due 579
529 to high organic matter production. In this last case, 580
530 Bacon and Anderson (1982) propose the “reversible 581
531 scavenging model”, in which the settling and regener- 582
532 ation of biogenic particles are important factors in 583
533 removing trace metals from surface waters in the ocean. 584
534 These factors are combined with surface water masses 585
535 present inside the bay. During period of study the water 586
536 column of Mejillones bay showed a mixture of Sub- 587
537 tropical Surface Water (STSW), Subantarctic Surface 588
538 Water (SASW), and Equatorial Subsurface Water 589
539 (ESSW), which are commonly observed offshore of 590
540 Mejillones Bay (Sobarzo and Figueroa, 2001). Consid- 591
541 ering that the bay have less than 100 m deep, it is 592
542 convenient to talk about the types of water associated to 593
543 these water masses. Their original properties are modi- 594
544 fied by a mixture of processes, and/or due to the atmo- 595
545 spheric conditions. The composition of these surface 596
546 water masses has a temporal and spatial variability that 597
547 is subject to seasonal atmospheric variations such as 598
548 solar radiation, wind intensity and a mixture of effects 599
549 caused by the bay’s dynamics and coastal upwelling 600
550 processes. The features of the water-mass core defined

in the oceanographic literature do not correspond to the 551
bay’s interior. However, it is possible to recognize these 552
traits in spite of the water mixture. In general terms, it 553
can be said that there is a predominance of the STSW at 554
the surface level. However, this level is also character- 555
ized by minimum salinities related to the minimum 556
salinity water levels (MSW) mentioned at Sobarzo and 557
Figueroa’s(2001) work corresponding to some type of 558
SASW. This situation should have an important 559
influence on surface distribution of trace metals and 560
can explain the high spatial and temporal variability 561
observed in most of the metals analyzed in this work. 562

563 Due to the fact that vertical mixing of water column 564
is an important characteristic in Mejillones bay, another 565
factor that should be considered is the influence of 566
microxic condition of subsurface waters (Valdés, 2004), 567
because all the metals analyzed in this work are redox- 568
sensitive (Morford and Emerson, 1999; Bostick et al., 569
2000; Brown et al., 2000, Crusius and Thomson, 2000; 570
Hammond et al., 2000; Zheng et al., 2000, Morford 571
et al., 2001; Russell and Morford, 2001). In the shallow 572
zone (between 0 m and 45 m water depth) sediments are 573
deposited under oxic condition. In this zone Ni, V, Cd 574
and Mo are resolubilized during the diagenetic process 575
remaining in solution, while Mn and Fe precipitate to 576
the sediments. Hatje et al. (2001) suggest that Fe is 577
oxidized faster than Mn in natural systems, which 578
indicates that Mn could be transported further away 579
from coast before precipitating in the bottom sediments. 580
In the deeper zone of Mejillones bay prevails a 581
microxic/anoxic sediment–water interface condition 582
which promotes precipitation of Ni, Cd, Mo and V. Fe 583
and Mn remain in solution and are resolubilized during 584
the diagenetic process.

585 The combination of these mechanisms and temporal 586
587 variation could explain the complex behavior of these 588
589 metals in this upwelling system of northern Chile. 590

4. Conclusions 588

589 The temporal distribution of some metals in the 590
591 coastal waters of the northern Humboldt Current 592
(Mejillones Bay, Chile) has been described in relation 593
to physical–chemical and biological processes. 594

595 In Mejillones Bay, nutrients are present year-round in 596
597 surface waters and are not a limiting factor for high 598
599 primary productivity. Radiation could be a limiting 600
factor explaining the opposing trend of nutrients and
chlorophyll; low radiation in autumn and winter cor-
responds to high nutrients and low chlorophyll, whereas
high radiation in spring and summer corresponds to low
nutrients and high chlorophyll.

The surface concentration of metals does not appear to be explained by anthropogenic inputs (at least not during the year of this work), and variability observed in this study appears to be natural. The lack of correlation between physical–chemical parameters and metals could indicate a more complex combination of factors acting on surface concentrations. It is possible, also, that the temporal variability of metals corresponds to a time scale other than those considered in this study. However, although not significant, a more intense variability in metal concentrations was observed at the pelagic station, which is under the direct influence of the Punta Angamos upwelling cell located off Mejillones Bay.

A two-box conceptual model is proposed to explain the variation of metals in surface waters of this coastal ecosystem. In the offshore box, surface circulation and deep water upwelling transport water masses with different metal contents and different temporal frequencies into the bay. In the inside the bay box, surface metal concentrations result from very low lithogenic input, ingestion by phytoplankton, desorption, and scavenging due to high organic matter production. The development of upwelling events in this area during the sampling period had different effects on each metals and sampling point. Moreover, the mixture of water masses and the Oxygen Minimum Zone which characterize the Mejillones bay should have an important influence on surface distribution of trace metals and can explain the high temporal variability observed in most of the metals analyzed in this work.

5. Uncited References

- Abe and Matsumaga, 1988
- Martín and Knauer, 1984
- Segovia-Zavala et al., 1998
- Takesue and Van Geen, 2002

Acknowledgements

This study was supported by grant PEI 1340-02, from the Universidad de Antofagasta, Chile, and UR 055 PALEOTROPIQUE, IRD, France. Support was also provided by IRD through the JEAI (Jeune Equipe Associée à l'IRD) program. We thank Freddy Rabasté and Keyla Majluf for their help in sampling and analytical support in nutrient and chlorophyll determinations. Special thanks to anonymous reviewers for their comments and significantly improving the manuscript.

References

- Abe, K., 2001. Cd in the western equatorial Pacific. *Marine Chemistry* 74, 197–211.
- Abe, K., Matsumaga, K., 1988. Mechanism controlling Cd and PO₄ concentration in Funka Bay, Japan. *Marine Chemistry* 23, 145–152.
- Adelson, J., Helz, G., Miller, C., 2001. Reconstructing the rise of recent coastal anoxia; molybdenum in Chesapeake bay sediments. *Geochimica et Cosmochimica Acta* 65 (2), 237–252.
- Bacon, M., Anderson, R., 1982. Distribution of thorium isotopes between dissolved and particulate forms in the deep sea. *Journal of Geophysical Research* 83, 2045–2056.
- Bostick, B., Fendorf, S., Fendorf, M., 2000. Disulfide disproportionation and CdS formation upon cadmium sorption on FeS₂. *Geochimica et Cosmochimica Acta* 64 (2), 247–255.
- Brown, J., Colling, A., Park, D., Phillips, J., Rothery D., Wright J., 1994. *Ocean Chemistry and Deep-Sea sediments*. The Open University/Pergamon Eds. Oxford, Great Britain: 133 pp.
- Brown, E., Callonec, L., German, Ch., 2000. Geochemical cycling of redox-sensitive metals in sediments from lake Malawi: a diagnostic paleotracer for episodic changes in mixing depth. *Geochimica et Cosmochimica Acta* 64 (20), 3515–3523.
- Bruland, K., 1983. Trace elements in sea-water. In: Riley, J.P., Chester, R. (Eds.), *Chemical Oceanography*, vol. 8. Academic Press, London, pp. 157–220.
- Calvert, S., Pedersen, T., 1993. Geochemistry of recent oxic and anoxic marine sediments: implications for the geological records. *Marine Geology* 11, 67–88.
- Cotté M. 1997. Origines et comportement des métaux dissous dans les eaux de la marge Atlantique Européenne. PhD Thesis, University of Paris VI, 210 pp.
- Cotte-Krief, M., Guieu, C., Thomas, A., Martin, J., 2000. Sources of Cd, Cu, Ni and Zn in the Portuguese coastal waters. *Marine Chemistry* 71, 199–214.
- Crusius, J., Thomson, J., 2000. Comparative behavior of authigenic Re, U, and Mo during reoxidation and subsequent long-term burial in marine sediments. *Geochimica et Cosmochimica Acta* 64 (13), 2233–2242.
- Cullen, J., Sherrell, R., 1999. Techniques for determination of trace metals in small samples of size-fractionated particulate matter: phytoplankton metals off central California. *Marine Geology* 67, 233–247.
- Dean, W., Gardner, J., Piper, D., 1997. Inorganic geochemical indicators of glacial–interglacial changes in productivity and anoxia on the California continental margin. *Geochimica et Cosmochimica Acta* 61 (21), 4507–4518.
- Delgadillo-Hinojosa, F., Macías-Zamorano, J., Segovia-Zavala, J., Torres-Valdés, S., 2001. Cadmium enrichment in the Gulf of California. *Marine Chemistry* 75, 109–122.
- Des Combes, H., Caulet, J., Tribovillard, N., 1999. Pelagic productivity changes in the equatorial area of the northwest Indian Ocean during the last 400,000 years. *Marine Geology* 15, 27–55.
- Escribano, R., Rodriguez, L., Iribarren, C., 1998. Temperature-dependent development and growth of *Calanus chilensis* Brodsky from northern Chile. *Journal of Experimental Marine Biology and Ecology* 229, 19–34.
- Escribano, R., Rosales, S., Blanco, J., 2004. Understanding upwelling circulation off Antofagasta (northern Chile): a three-dimensional numerical-modeling approach. *Continental Shelf Research* 24, 37–53.

- Fahrbach, E., Brockman, C., Lostanau, N., Urquizo, W., 1980. The northern peruvian upwelling system during the ESACAN experiment. In: Richards, F. (Ed.), Coastal Upwelling, Coastal and Estuarine Science. AGU, Washington D.C., pp. 134–145.
- Gonzalez, H., Daneri, G., Figueroa, D., Iriarte, J., Lefevre, N., Pizarro, G., Quiñónez, R., Sobarzo, M., Troncoso, A., 1998. Producción Primaria y su destino en la trama trófica pelágica y océano profundo e intercambio océano-atmósfera de CO₂ en la zona norte de la corriente de Humboldt (23° S): posibles efectos del evento El Niño, 1997–98 en Chile. *Revista Chilena de Historia Natural* 71, 429–458.
- Gonzalez, H., Ortiz, V., Sobarzo, M., 2000. The role of faecal material in the particulate organic carbon flux in the northern Humboldt Current, Chile (23° S), before and during the 1997–1998 El Niño. *Journal of Plankton Research* 22 (3), 499–529.
- Gonzalez-Davila, H., 1995. The role of phytoplankton cells on the control of heavy metals concentrations in seawater. *Marine Chemistry* 48, 215–236.
- Gonzalez, H., Giesecke, R., Vargas, C., Pavez, M., Iriarte, J., Santibáñez, P., Castro, L., Escribano, R., Pagés, F., 2004. Carbon cycling through the pelagic foodweb in the northern Humboldt Current off Chile (23° S). *Journal of Marine Science* 61, 572–584.
- Grotti, M., Soggia, F., Abelloschi, M., Rivaro, P., Magi, E., Frache, R., 2001. Temporal distribution of trace metals in Antarctic coastal waters. *Marine Chemistry* 76, 189–209.
- Hammond, D., Mcmanus, J., Berenson, W., Meredith, Ch., Klinkhammer, G., Coale, K., 2000. Diagenetic fractionation of Ge and Si in reducing sediments: the missing Ge sink and a possible mechanism to cause glacial/interglacial variations in oceanic Ge/Si. *Geochimica et Cosmochimica Acta* 64 (14), 2453–2465.
- Hatje, V., Birch, G., Hill, M., 2001. Spatial variability of particulate trace metals in Port Jackson Estuary, Australia. *Estuarine, Coastal and Shelf Science* 53, 62–77.
- Holm-Hansen, O., Lorenzen, C., Holmes, R., Strickland, J., 1965. Fluorometric determination of chlorophyll. *Journal du Conseil - Conseil Permanent International pour l'exploration de la Mer* 30, 3.15.
- Hutchins, D., Bruland, K., 1998. Iron-limited diatom growth and Si:N uptake ratios in a coastal upwelling regime. *Nature* 393, 561–564.
- Iriarte, J., Gonzalez, H., 2004. Phytoplankton size structure during and after the 1997/1998 El Niño in coastal upwelling area of the northern Humboldt Current System. *Marine Ecology Progress Series* 269, 83–90.
- Iriarte, J., Pizarro, G., Troncoso, V., Sobarzo, M., 2000. Primary production and biomass of size-fractionated phytoplankton of Antofagasta, Chile (23–24° S) during pre-El Niño and El Niño 1997. *Journal of Marine Systems* 26, 37–51.
- Johnson, K., Chavez, F., Friedrich, G., 1999. Continental-shelf sediments as a primary source of iron for coastal phytoplankton. *Nature* 398, 697–700.
- Kremling, K., 1985. The distribution of cadmium, copper, nickel, manganese, and aluminum in surface waters of the open Atlantic and European shelf area. *Deep-Sea Research* 32 (5), 531–555.
- Kremling, K., Hydes, D., 1988. Summer distribution of dissolved Al, Cd, Cu, Mn, and Ni in the surface waters around the British Isles. *Continental Shelf Research* 8, 89–105.
- Kremling, K., Pohl, C., 1989. Studies on the spatial and seasonal variability of dissolved cadmium, copper and nickel in the North-east Atlantic surface waters. *Marine Chemistry* 27, 43–60.
- Le Gall, A., Hydes, D., Statham, P., Morley, H., Hunt, C., 1999. Processes influencing distribution and concentration of Cd, Cu, Mn and Ni at the North West European shelf breaks. *Marine Chemistry* 68, 97–115.
- Libes, S., 1992. An Introduction to Marine Biogeochemistry. John Wiley & Sons, Inc., pp. 15–190.
- Marín, V., Olivares, G., 1999. Estacionalidad de la productividad primaria en bahía Mejillones del Sur (Chile): una aproximación proceso-funcional. *Revista Chilena de Historia Natural* 72, 629–641.
- Marín, V., Escribano, R., Delgado, L., Olivares, G., Hidalgo, P., 2001. Nearshore circulation in a coastal upwelling site off the northern Humboldt Current System. *Continental Shelf Research* 21, 1317–1329.
- Marín, V., Delgado, L., Escribano, R., 2003. Upwelling shadows at Mejillones bay (northern Chilean coast): a remote sensing *in situ* análisis. *Investigaciones Marinas* 31 (2), 47–55.
- Martín, J., Knauer, G., 1984. Manganese transport through oxygen minima. *Earth and Planetary Science Letters* 67, 35–47.
- Martín, J., Fitzwater, S., 1988. Iron deficiency limits phytoplankton growth in the north-east Pacific subarctic. *Nature* 331, 341–343.
- Martín, J., Thomas, A., 1994. The global insignificance of telluric input of dissolved trace metals (Cd, Cu, Ni and Zinc) to ocean margin. *Marine Chemistry* 46, 165–178.
- Morford, J., Emerson, S., 1999. The geochemistry of redox-sensitive trace metals in sediments. *Geochimica et Cosmochimica Acta* 63 (11/12), 1735–1750.
- Morford, J., Russell, A., Emerson, S., 2001. Trace metal evidence for changes in the redox environment associated with the transition from terrigenous clay to diatomaceous sediment, Saanich Inlet, BC. *Marine Geology* 174, 355–369.
- Morley, N., Statham, P., Burton, J., 1993. Dissolved trace metals in the southwestern Indian Ocean. *Deep-Sea Research* 40 (5), 800–801.
- Muller, F., Tranter, M., Balls, P., 1994. Distribution and transport of chemical constituents in the Clyde Estuary. *Estuarine, Coastal and Shelf Science* 39, 105–126.
- Nameroff, T., Balistrieri, L., Murray, W., 2002. Suboxic trace metals geochemistry in the eastern tropical North Pacific. *Geochimica et Cosmochimica Acta* 66 (7), 1139–1158.
- Ortlieb, L., Escribano, R., Follegati, R., Zúñiga, O., Kong, I., Rodríguez, L., Valdes, J., Guzmán, N., Iratchet, P., 2000. Ocean-climatic changes during the last 2,000 years in a hypoxic marine environment of Northern Chile (23° S). *Revista Chilena de Historia Natural* 73, 221–242.
- Parsons, T., Maita, Y., Lalli, C., 1984. A Manual of Chemical and Biological Methods Seawaters Analysis. Pergamon Press, Oxford, United Kingdom, p. 173.
- Pizarro, G., Iriarte, J., Montecinos, V., 2002. Mesoscale primary production and bio-optical variability off Antofagasta (23°–24°S) during the transition to El Niño 1997–1998. *Revista Chilena de Historia Natural* 75, 201–215.
- Rodríguez, L., Marín, V., Farias, M., Oyarce, E., 1991. Identification of an upwelling zone by remote sensing and *in situ* measurement. *Mejillones del Sur Bay (Antofagasta-Chile)*. *Scientia Marina* 55 (3), 467–473.
- Román, D., Rivera, L., Morales, T., Ávila, J., Cortés, P., 2003. Determination of trace elements in environmental and biological samples using improved simple introduction in flame absorption spectrometry (HHPN-AAS; HHPN-FF-AAS). *International Journal of Environmental Analytical Chemistry* 83, 327–341.
- Russell, A., Morford, J., 2001. The behavior of redox-sensitive metals across a laminates-massive-laminated transition in Saanich Inlet, British Columbia. *Marine Geology* 174, 341–354.
- Saager, P., De Baarm, H., Howland, R., 1992. Cd, Zn, Ni and Cu in the Indian Ocean. *Deep-Sea Research* 39 (1), 9–35.

- 834 Schenau, S., Reichart, G., De Lange, G., 2002. Oxygen minimum zone
835 controlled Mn redistribution in Arabian Sea sediments during the
836 late Quaternary. *Paleoceanography* 17 (4), 10–1–10–12.
- 837 Segovia-Zavala, J., Delgadillo-Hinojosa, F., Alvarez-Borrego, S.,
838 1998. Cadmium in the coastal upwelling area adjacent to the
839 California–Mexico border. *Estuarine, Coastal and Shelf Science*
840 46, 475–781.
- 841 Sobarzo, M., Figueroa, D., 2001. The physical structure of a cold
842 filament in a Chilean upwelling zone (Península de Mejillones,
843 Chile, 23° S). *Deep-Sea Research* 48, 2699–2726.
- 844 Strub, P., Mesias, J., Montecinos, V., Rutllant, J., Marchant, S., 1998.
845 Coastal oceanic circulation off western South America. In:
846 Robinson, A., Birk, K. (Eds.), *The Sea*. John Wiley, New York,
847 USA, pp. 273–314.
- 848 Takesue, R., Van Geen, A., 2002. Nearshore circulation during
849 upwelling inferred from the distribution of dissolved cadmium off
850 the Oregon coast. *Limnology and Oceanography* 47, 176–185.
- 851 Takesue, R., Van Geen, A., Carriquiry, J., Ortiz, E., Godínez-Orta, L.,
852 Granados, I., Saldivar, M., Ortlieb, L., Escribano, R., Guzmán, N.,
853 Castilla, J., Varas, M., Salamanca, M., Figueroa, C., 2004. Influence
854 of coastal upwelling and El Niño–Southern Oscillation on nearshore
855 water along Baja California and Chile: shore-based monitoring
856 during 1997–2000. *Journal of Geophysical Research* 109, 1–14.
- 857 Ulloa, O., Escribano, R., Hormazabal, S., Quiñónez, R., Gonzalez, R.,
858 Ramos, R., 2001. Evolution and biological effects of the 1997–98
859 El Niño in the upwelling ecosystem off northern Chile.
860 *Geophysical Research Letters* 28, 1591–1594.
- 861 Valdés, J., 2004. Evaluación de metales redox-sensitivos como proxies
862 de paleoxigenación en un ambiente marino hipóxico del norte de
863 Chile. *Revista Chilena de Historia Natural* 77, 121–138.
- 864 Valdés, J., Ortlieb, L., 2001. Paleoxigenación subsuperficial de la
865 columna de agua en la bahía Mejillones del Sur (23° S): Indicadores
866 geoquímicos en testigos de sedimento marino. *Investigaciones
867 Marinas* 29 (1), 25–35.
- Valdés, J., López, L., Lomónaco, S., Ortlieb, L., 2000. Condiciones 868
paleoambientales de sedimentación y preservación de materia 869
orgánica en Bahía Mejillones del Sur (23° S), Chile. *Biología 870
Marina y Oceanografía. Revista de Biología Marina y Oceano- 871
grafía* 35 (2), 169–180. 872
- Valdés, J., Ortlieb, L., Sifeddine, A., 2003. Variaciones del sistema de 873
surgencia de Punta Angamos (23° S) y la Zona de Mínimo 874
Oxígeno durante el pasado reciente. Una aproximación desde el 875
registro sedimentario de la bahía Mejillones del Sur. *Revista 876
Chilena de Historia Natural* 76, 347–362. 877
- Valdés, J., Sifeddine, A., Lallier-Verges, E., Ortlieb, L., 2004. 878
Petrographic and geochemical study of organic matter in surficial 879
sediments from an upwelling system (Mejillones del Sur bay, 880
northern Chile). *Organic Geochemistry* 35 (5), 881–894. 881
- Valdés, J., Vargas, G., Sifeddine, A., Ortlieb, L., Guíñez, M., 2005. 882
Distribution and enrichment evaluation of heavy metals in 883
Mejillones Bay (23° S), northern Chile: geochemical and statistical 884
approach. *Marine Pollution Bulletin* 50, 1558–1568. 885
- Van Geen, A., Husby, D., 1996. Cadmium in the California Current 886
System: tracer of the past and present upwelling. *Journal of Geo- 887
physical Research* 101, 3489–3507. 888
- Van Geen, A., Luoma, S., Fuller, C., Anima, R., Clifton, E., Trumbore, 889
S., 1992. Evidence from Cd/Ca ratio in foraminifera for greater 890
upwelling off California 4,000 years ago. *Nature* 358, 54–56. 891
- Vargas, G., Ortlieb, L., Pichon, J., Bertaux, J., Pujos, M., 2004. 892
Sedimentary facies and high resolution primary production infer- 893
ences from laminated diatomaceous sediments off northern Chile 894
(23° S). *Marine Geology* 211, 79–99. 895
- Zheng, Y., Anderson, R., Van Geen, A., Kuwabara, J., 2000. 896
Authigenic molybdenum formation in marine sediments: a link 897
to porewater sulfide in the Santa Barbara basin. *Geochimica et 898
Cosmochimica Acta* 64, 4165–4178. 899